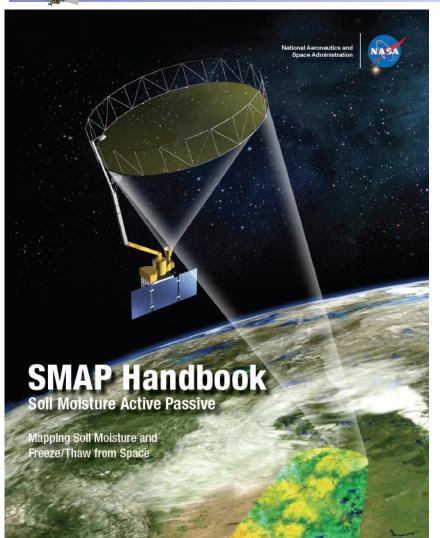




#### **SMAP Mission Concept**





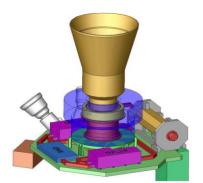
L-band unfocused SAR and radiometer system, offset-fed 6 m light-weight deployable mesh reflector. Shared feed for

- > 1.2 GHz Radar 1-3 km (30% nadir gap)
  HH, VV and HV
- > 1.4 GHz Radiometer at 40 km (-3 dB) H, V, 3<sup>rd</sup> and 4<sup>th</sup> Stokes

Conical scan, fixed incidence angle at 40°

Contiguous 1000 km swath 2-3 days revisit

Sun-synchronous 6am/6pm orbit (680 km)



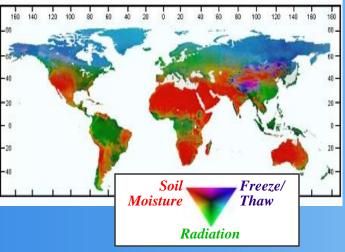
Electronic Version at http://smap.jpl.nasa.gov/Imperative/
Print Version Available (182 Pages): smap\_science@jpl.nasa.gov

# **Applications Returns**

#### SMAP Science and Application Returns



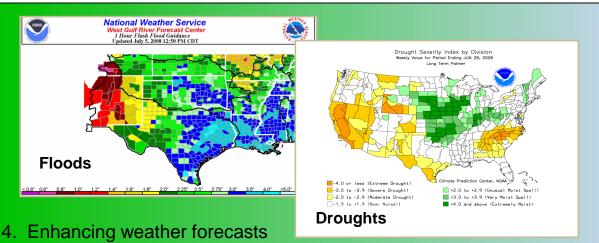
Soil Moisture <u>Links</u> the Global Land Water, Energy, and Carbon Cycles



- Estimating global surface water and energy fluxes
- Quantifying net carbon flux in boreal landscapes
- 3. Reducing uncertainty of climate model projections



L-band (~21 cm; All-Weather; Canopy Penetration; Sensing Depth)



5. Improving flood prediction and drought monitoring

6m conically scanning (14 rpm) antenna for 1000 km swath

Global coverage every 2-3 days



#### Core Science Objective

SECTION F



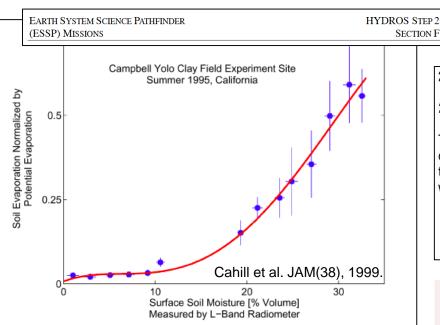


Figure F.1-2. A ground-based L-band radiometer is used to make the soil moisture field measurements to estimate the surface control on evaporation (fitted red line). Global HY-DROS soil moisture measurements, together with meteorological and hydrological data, will allow for the first time a quantification of influential processes such as this across diverse climatic and seasonal regimes.

SMAP L1 Science Requirements and Mission Success Criteria

#### 2.0 SCIENCE DEFINITION

#### 2.1 BASELINE SCIENCE OBJECTIVES

The SMAP Project will implement a spaceborne earth observation mission designed to collect measurements of surface soil moisture and freeze/thaw state. together termed the hydrosphere state. SMAP hydrosphere state measurements will yield a critical data set that will enable science and applications users to:

- · Understand processes that link the terrestrial water, energy and carbon cvcles:
- Estimate global water and energy fluxes at the land surface:

Multiple approaches being pursued and coordinated (R. Koster, G. Salvucci, J. Kimball, D. Entekhabi and others).

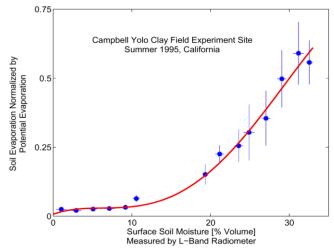
Estimate with least reliance on models and parameterizations (i.e., be observations-driven).

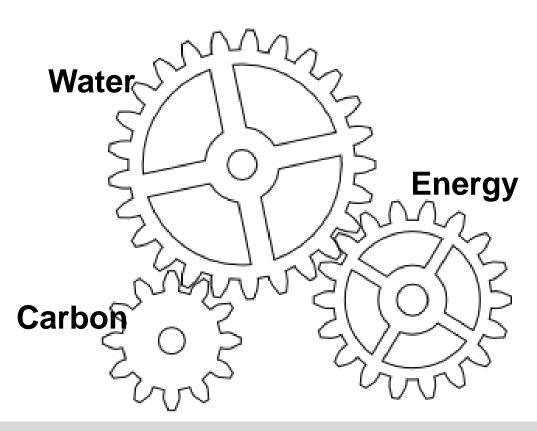
Relate function to vegetation type, seasonal climate and soil texture.



# Sources of Earth System Model Discrepancies and Uncertainties







Cycles are Gears That Turn Together

But The Closure Relations Linking Them are Parameterized

This is a Major Source of Uncertainty in Models

# Demonstration Model: Role of Water-Energy-Carbon Balance Closure Function

$$\frac{d \text{ Soil Water}}{dt} = \text{ Precipitation - Evaporation}$$

$$\frac{d \text{ Temperature}}{dt} = \text{ Incoming Radiation - Turbulent Heat Flux}$$

$$\frac{d \text{ Carbon}}{dt} = \text{ Assimilation - Respiration}$$

$$\frac{d \text{ Carbon}}{dt} = \text{ Assimilation - Respiration}$$

$$\frac{d \text{ Soil Water}}{dt} = P - E(S, T)$$

$$\frac{d \text{ Temperature}}{dt} = R^{-} - E_{s}ST^{4} - L \times E(S, T) - H$$

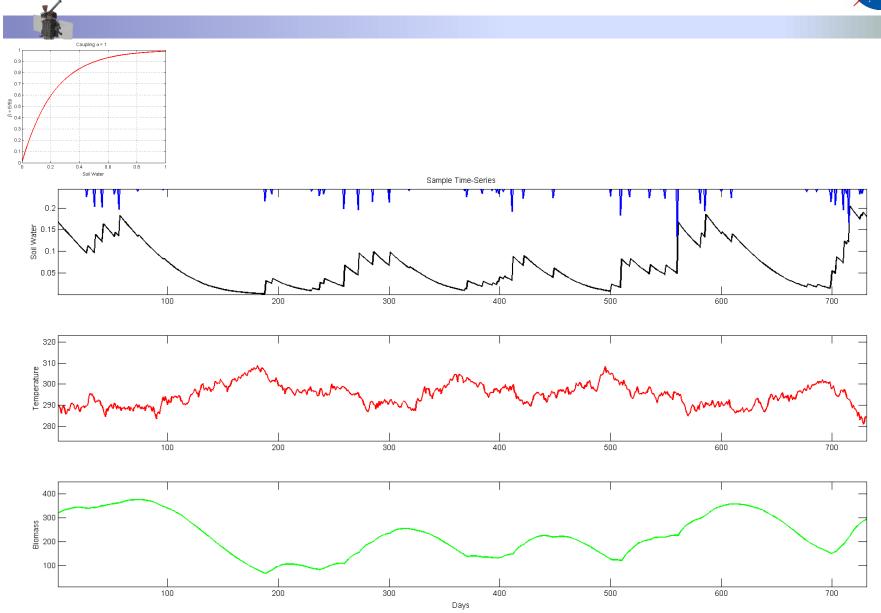
$$\frac{d \text{ Carbon}}{dt} = C \times E(S, T) - \frac{B}{t}$$

All coupled together through:  $E(S,T) = b(S) \times E_p$ 

 $\beta(S)$  is hence the <u>closure function</u> between the water, energy and carbon budgets.

#### Three-Variable Demonstration Model

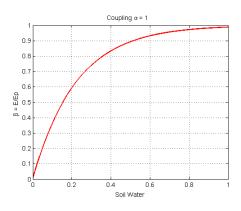




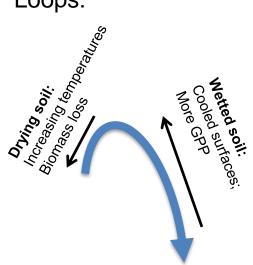


# Co-Variations of States: Dependence on Closure Function

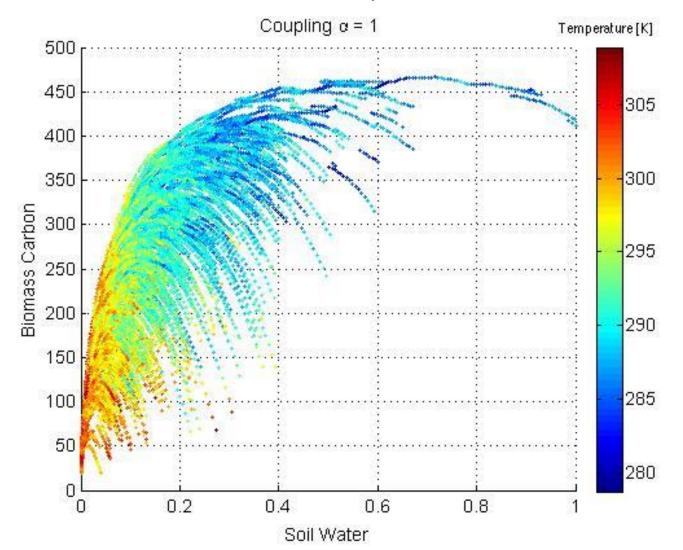




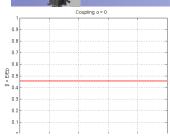
## Inter-Storm Loops:



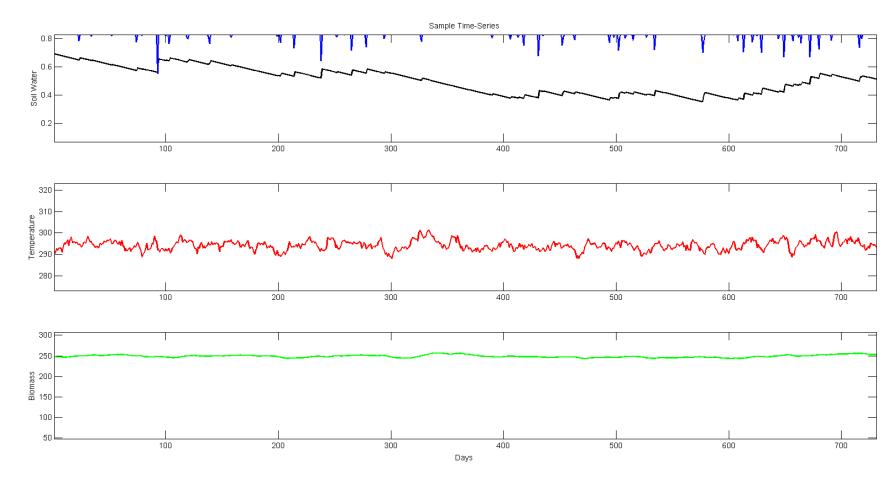
## Sample state-space diagram for the demonstration three-equation model.







Without the closure function (see left) co-variations muted in the three variables.





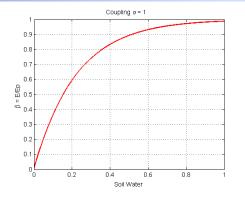
#### Cycles Interrupted

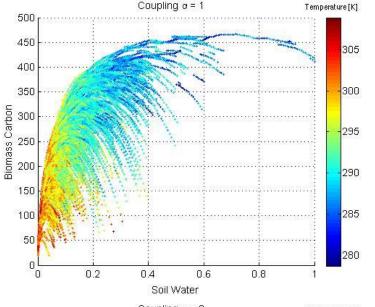


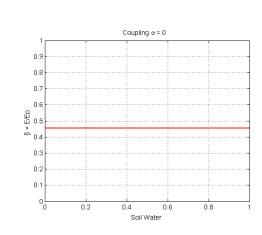
Sample state-space diagram for the demonstration three-equation model.

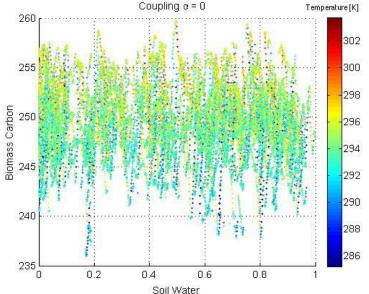
Without closure, covariations disappear.

But this is extreme contrast case.











# Parameterized Closure Functions But Without Strong Evidence



#### **NOAH**

model grid cell and

$$\beta = \left(\frac{\Theta_1 - \Theta_w}{\Theta_{\text{ref}} - \Theta_w}\right)^f \tag{7}$$

represents a normalized soil moisture availability term where  $\Theta_w$  is the wilting point and  $\Theta_{ref}$  is the field capac-

#### CLM

functional type and the soil water potential of each soil layer

$$\beta_t = \sum_i w_i r_i \ge 1 \times 10^{-10} \tag{8.10}$$

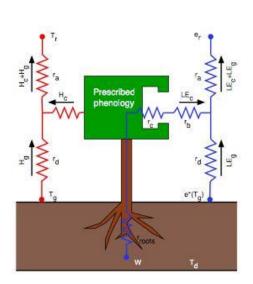
where  $w_i$  is a soil dryness or plant wilting factor for layer i, and  $r_i$  is the fraction of roots in layer i.

The plant wilting factor w, is

$$v_{i} = \left\{ \frac{\psi_{\text{max}} - \psi_{i}}{\psi_{\text{max}} + \psi_{\text{sat}, i}} \qquad \text{for } T_{i} > T_{f} \right\}$$

$$(8.11)$$

$$\beta = \begin{cases} \frac{1}{4} \left[ 1 - \cos\left(\frac{\theta_1}{\theta_{fc}}\pi\right) \right]^2 & \theta_1 < \theta_{fc} \\ 1 & \theta_1 \ge \theta_{fc} & \text{or} \quad q_{air} > \alpha q_{sat}(T_g) \end{cases}, \tag{5}$$

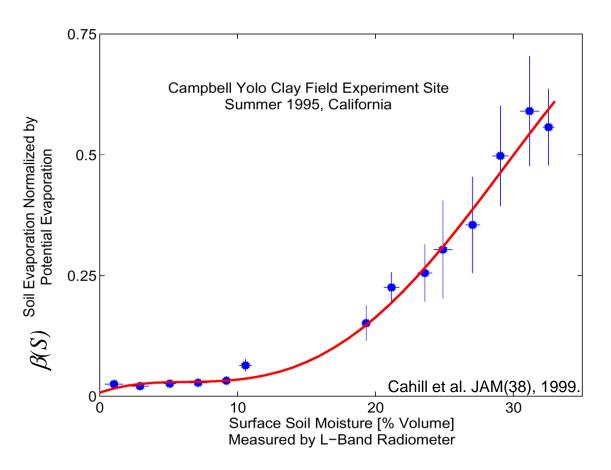


R. Stöckli and P. L. Vidale (ETH)

#### How well have we measured/estimated?



Only measured at a few flux tower sites. But valid for the footprint of flux tower (does not scale to a global model grid area) and only known for a few landscape types.



To estimate this closure function, independent observations of

soil moisture state and

evaporation flux

are required. Globally.



#### Evaporation



100

80

60

40

20

**Emergent relation between surface vapor conductance** and relative humidity profiles yields evaporation rates from weather data

Guido D. Salvucci<sup>a,1</sup> and Pierre Gentine<sup>b</sup>

<sup>a</sup>Department of Earth and Environment, Boston University, Boston, MA 02215; and <sup>b</sup>Department of Earth and Environmental Engineering, Columbia

Water Resources Research

RESEARCH ARTICLE

Evapotranspiration based on equilibrated relative humidity

(ETRHEQ): Evaluation over the continental U.S.

10.1002/2014WR016072

Surface conductance is estimated

Angela J. Rigden<sup>1</sup> and Guido D. Salvucci<sup>1</sup>

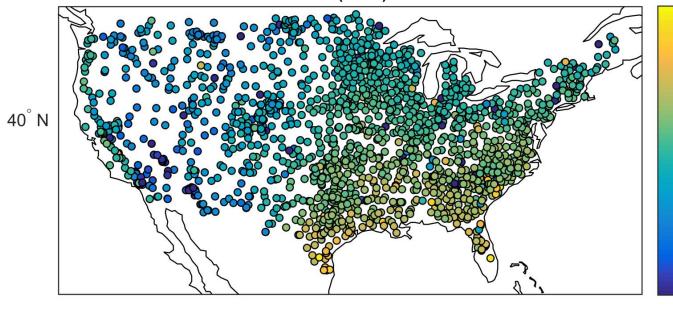
<sup>1</sup>Department of Earth and Environment, Boston University, Boston, Massachusetts, USA



Guido Salvucci Pierre Gentine (Boston Univ.) (Columbia Univ.)



LE (W/M<sup>2</sup>)

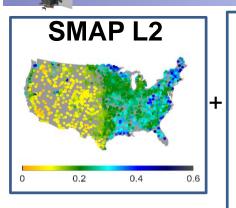


April 1, 2015 to March 30, 2016

> 120° W 100° W

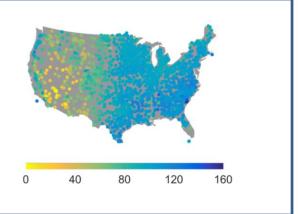
## Coupling of Terrestrial Water, Energy and Carbon Cycles

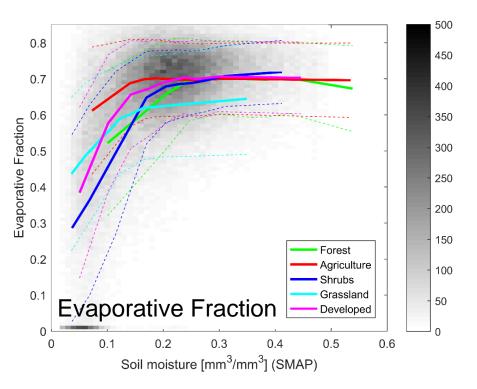


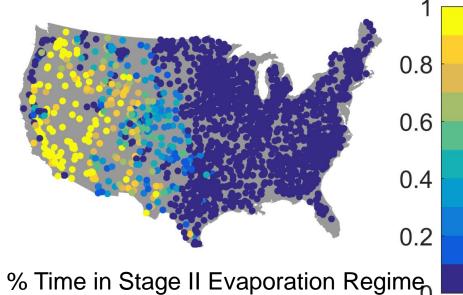


## ETRHEQ (Evaporation Rate)

(Salvucci & Gentine, PNAS 2013)





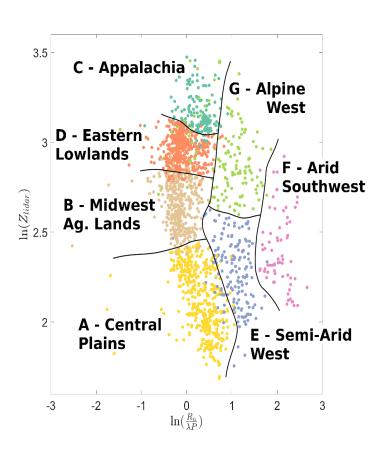


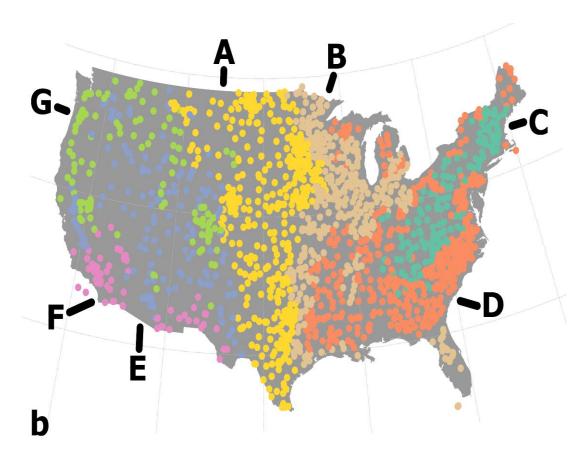
Gianotti, Rigden, Salvucci and Entekhabi (2017)



## **Ecohydrological Regions**









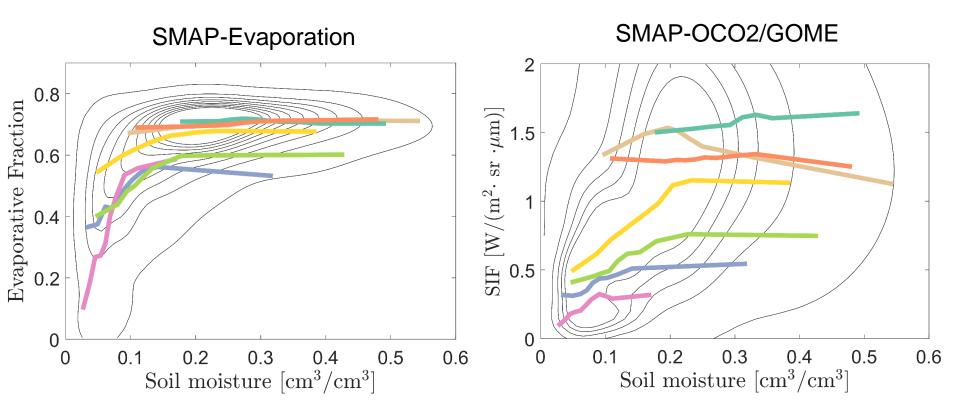
Central Plains

Appalachia

Midwest Agricultural Lands E − Semi-arid West

#### Soil Moisture Controls on Water and Carbon Cycles





F – Arid Southwest

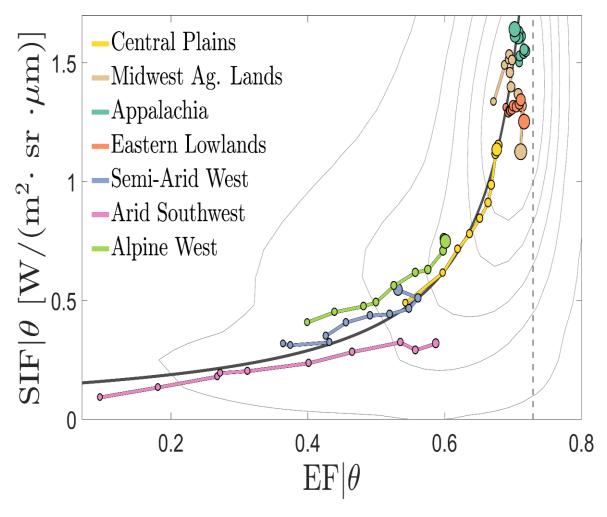
D – Eastern Lowland Forests G – Alpine West



#### $SIF|\theta \text{ vs } EF|\theta$



Gianotti, Rigden, Salvucci and Entekhabi (2017)



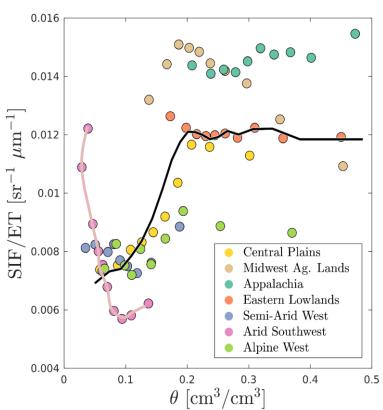
- Connects EF vs θ and SIF vs θ curves
- AIC\prefer a single relationship over independent relationships by region
- Convex shape implies landscape-level Water Use Efficiency increases with EF (opposite of physiological WUE)



#### Landscape-Level "Water Use Efficiency (WUE)"

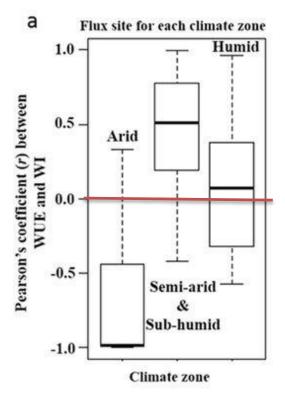


#### SIF/ET is WUE at the landscape scale



Gianotti, Rigden, Salvucci and Entekhabi (2017)

Decreases with  $\theta$  for the arid regions, then increases with  $\theta$  for semi-arid regions and flat for humid regions (in agreement with Yang *et al.*<sup>1</sup>)



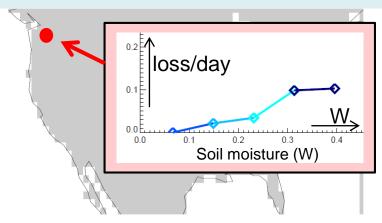
[1] Yang, Y. et al. Contrasting responses of water use efficiency to drought across global terrestrial ecosystems. *Nature*, **6**, 23284 (2016). [WI=Wetness Index]\_

#### SMAP Estimates of Surface Hydrology Loss Functions



#### R. Koster (NASA GSFC)

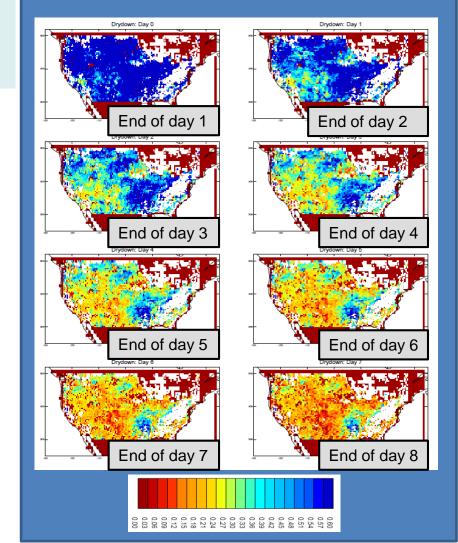
Spatial distributions of soil moisture loss function (loss from evaporation and drainage) can be extracted from collocated SMAP and precipitation measurements for summertime:



Derived loss functions are highly valuable:

- Improved precipitation estimation
- "Filling-in" of missing data; accurate forecasts of soil moisture several days into the future
- Characterization of regional differences in hydrological behavior (figure on right)

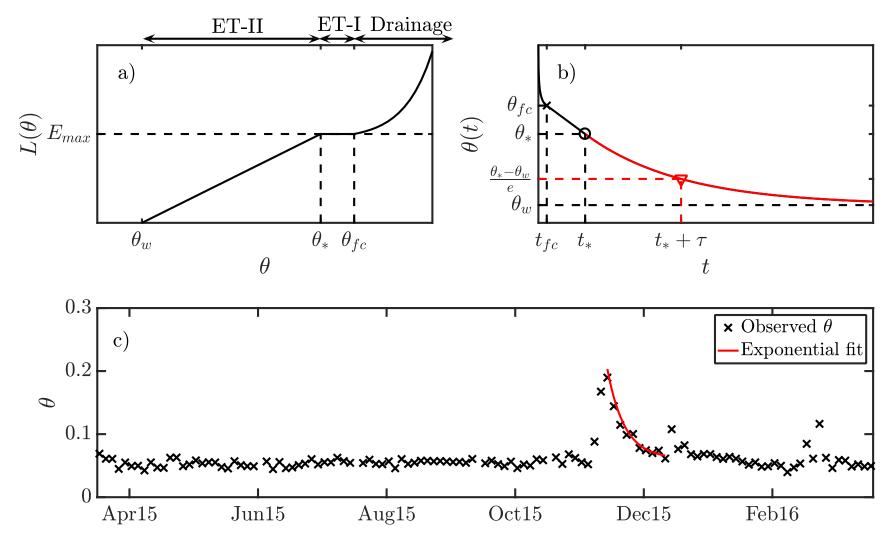
Loss function analysis: Begin with wet soil everywhere and examine *drydown* 





## Stages of Drydowns: Landscape Water Loss

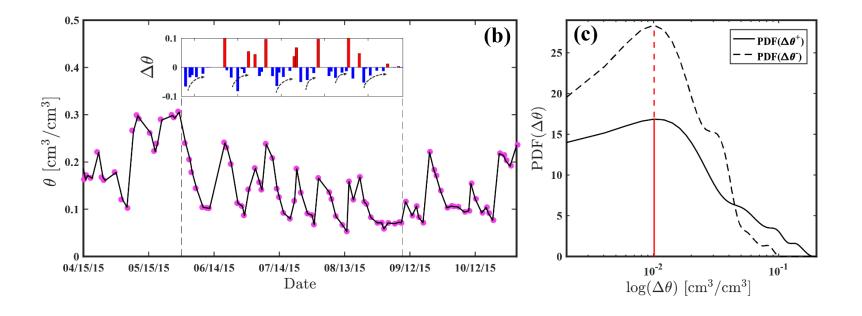






## Landscape Water Loss Function: Observation-Driven Estimates





$$\Delta \theta_{i_{+}} = \begin{cases} \Delta \theta_{i}, & \text{if } \Delta \theta_{i} > 0 \\ 0, & \text{otherwise} \end{cases}$$

$$\Delta \theta_{i_{-}} = \begin{cases} \Delta \theta_{i}, & \text{if } \Delta \theta_{i} < 0 \\ 0, & \text{otherwise} \end{cases}$$

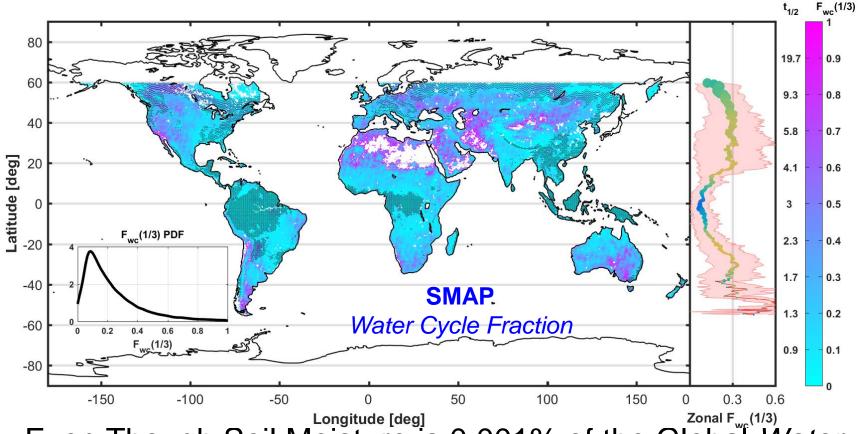


## Soil Moisture and the Terrestrial Water Cycle: Positive Increments



McColl, Alemohammad, Akbar, Konings, Yueh and Entekhabi, 2017: The global distribution and dynamics of surface soil moisture, *Nature-Geoscience*, 10(2).

$$\Delta\theta_{i_{+}} = \begin{cases} \Delta\theta_{i}, & \text{if } \Delta\theta_{i} > 0 \\ 0, & \text{otherwise} \end{cases}$$



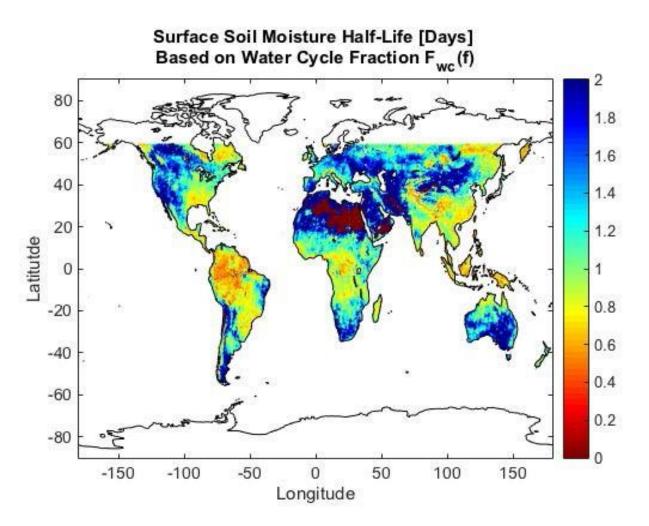
Even Though Soil Moisture is 0.001% of the Global Water Budget, it Captures About 20% of the Water Cycle



## Soil Moisture Memory



$$t_{1/2} = -\frac{f^{-1}}{\log_2(F_{wc}(f))}$$



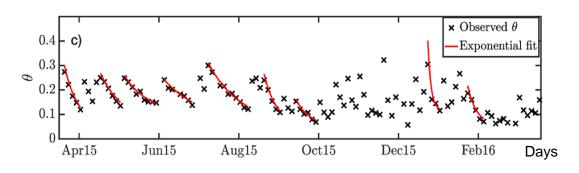


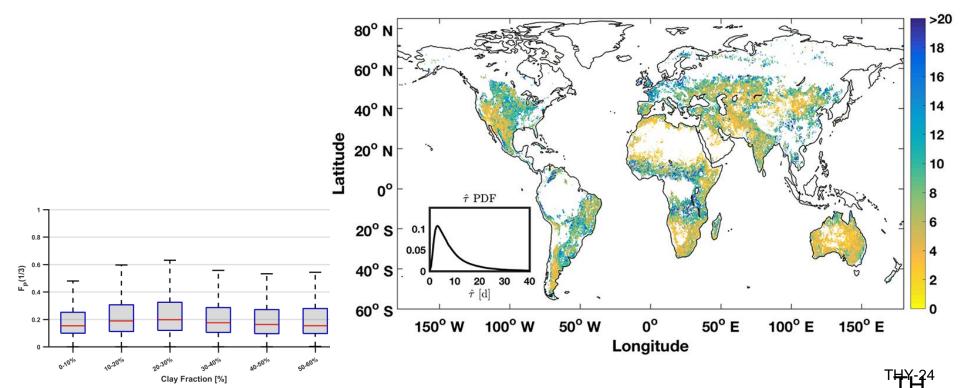
#### Soil Moisture Memory



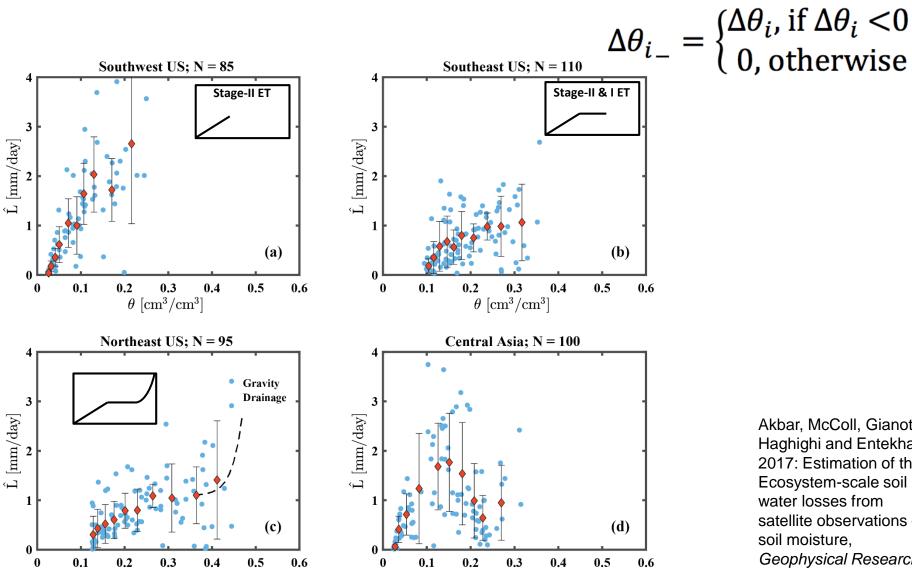
Memory of surface soil moisture can extend to days especially where land-atmosphere coupling is significant and forecast skill can be extended.

McColl et al., 2017b: Global characterization of surface soil moisture drydowns, 44, *Geophysical Research Letters*.





## Observation-Driven Estimates of Landscape Water Loss: **Negative Increments**



 $\theta \, [\mathrm{cm}^3/\mathrm{cm}^3]$ 

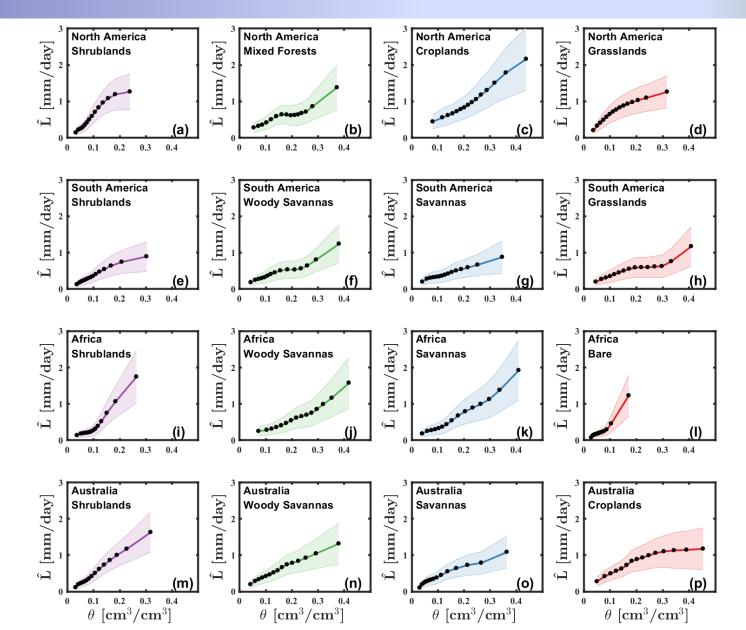
 $\theta \, [\mathrm{cm}^3/\mathrm{cm}^3]$ 

Akbar, McColl, Gianotti, Haghighi and Entekhabi, 2017: Estimation of the Ecosystem-scale soil water losses from satellite observations of Geophysical Research Letters



#### Loss Function for Different IGBP Classes

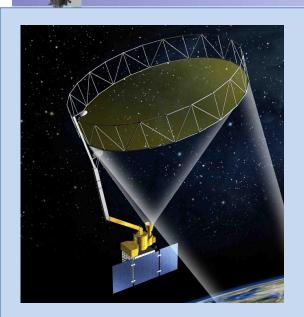






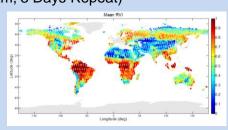
# Global Ecology in the SMAP, OCO-2, and ICESat-2 Era





Radar Vegetation Index (3 km, 4 Days Repeat)

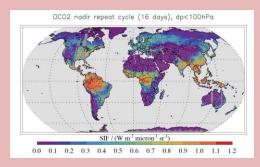
## Radiometer Vegetation µwave Opacity (40 km, 3 Days Repeat)



**Canopy Water Content**and Stress



## Solar-Induced Chlorophyll Fluorescence

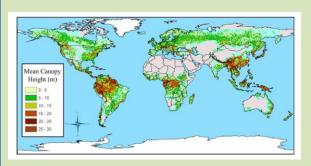


Frankenberg et al., 2014: Prospects for chlorophyll fluorescence remote sensing from the OCO-2, RSE 147, 1-12.

**GPP** 



#### Canopy Height



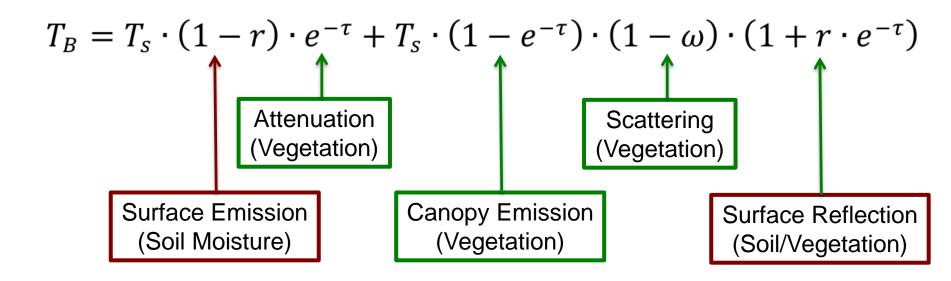
Abdalati et al., The ICESat-2 Laser Altimetry mission, Proc. IEEE, 98, 735-751.

Structure



#### Radiative Transfer





#### Beyond Baseline Algorithm:

Observations:  $TB_H$ ,  $TB_V$ 

Unknowns:  $\varepsilon_s$ ,  $\tau$ ,  $\omega$ 

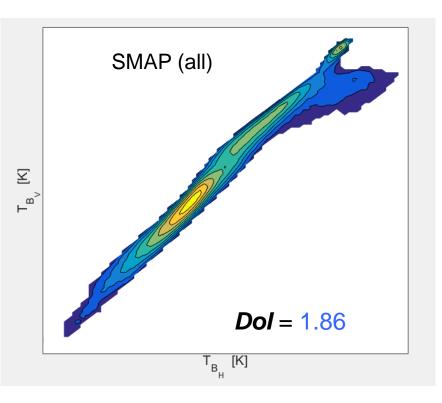


## Retrieving Vegetation Effects Without Relying on Optical Ancillary Data

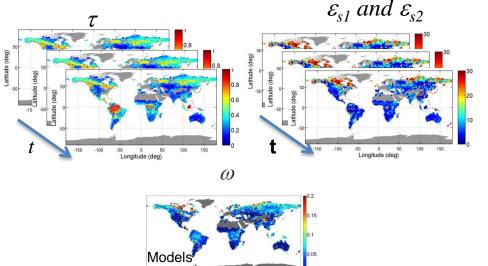


**Dol**: Konings, McColl, Piles, Entekhabi, 2015: How Many Parameters Can Be Maximally Estimated From a Set of Measurements? *GRSL* 12(5)

$$I(X;Y) = \int_{Y} \int_{X} p(x,y) \log \left( \frac{p(x,y)}{p(x)p(y)} \right) dx dy,$$



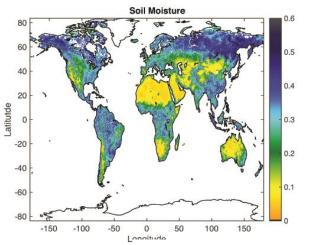
N = 2 overpasses (4 Measurements):  $2 \times DoI = 2 \times 1.86 = 3.7$ 



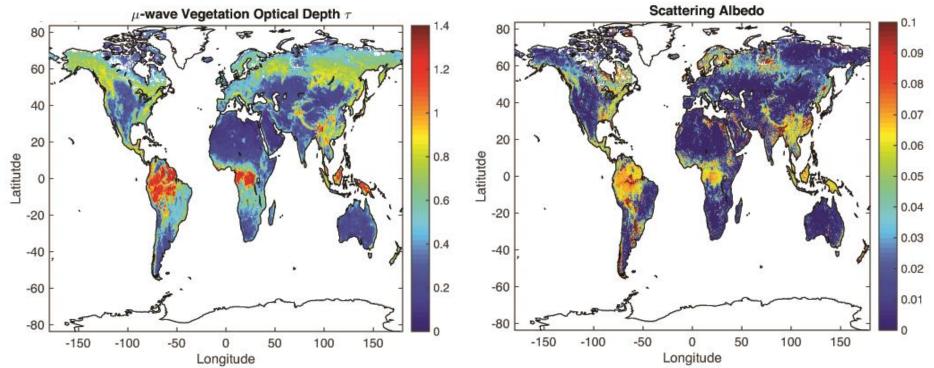


#### Climatology of Retrievals





Konings, Piles, Das and Entekhabi, 2017: L-Band vegetation optical depth and scattering albedo estimation from SMAP, *Remote* Sensing of the Environment

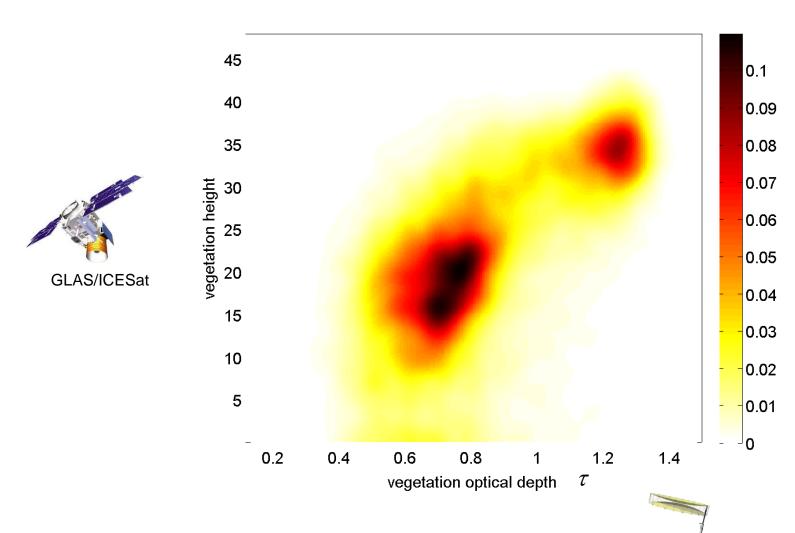




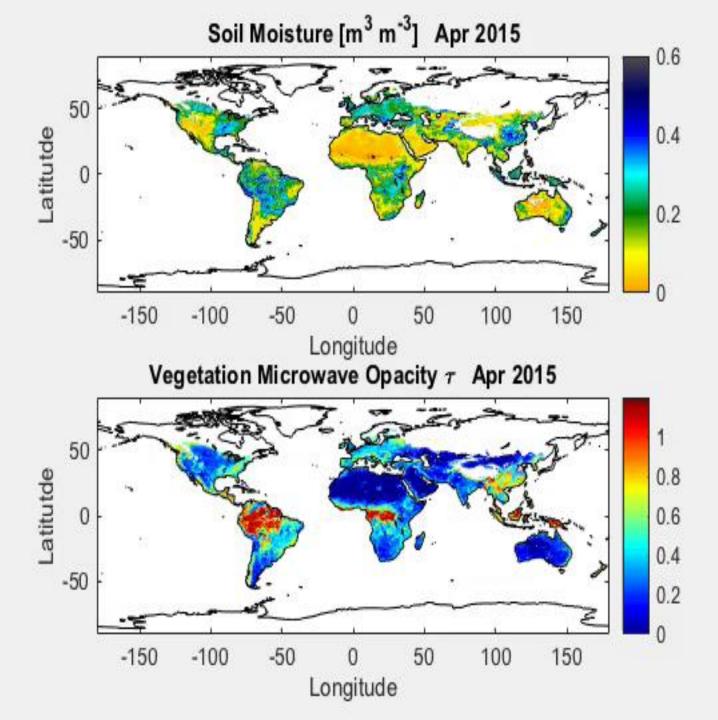
#### Microwave Interactions With Vegetation



**SMAP** 



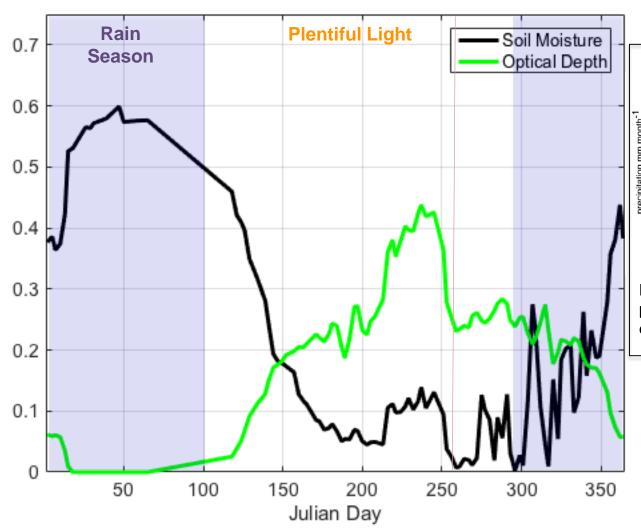
Rötzer, Montzka, Entekhabi, Konings, McColl, Piles, Vereecken, 2017: Relationship between vegetation optical depth and HV-backscatter from the Aquarius mission, *IEEE Transactions of Geoscience and Remote Sensing*.

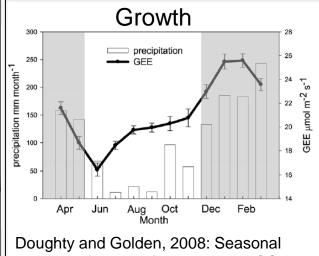




## Dry Tropical Forest Phenology







patterns of tropical forest LAI and CO2 exchange, *JGR-Biogeosciences*, vol. 113.

#### Flushing

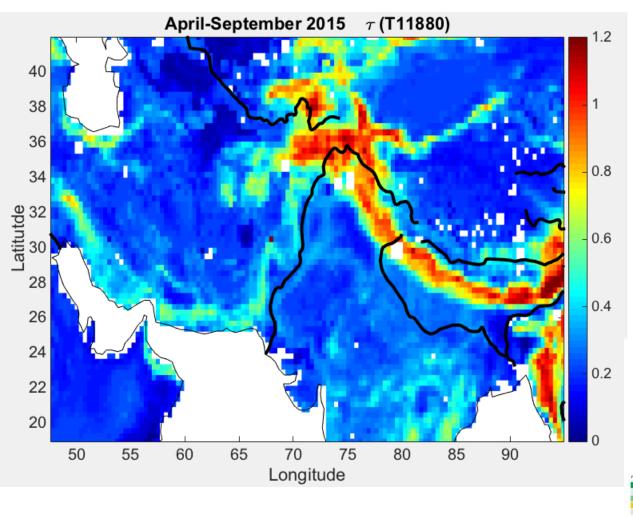
Rivera et al., 2002: Increasing day-length induces spring flushing of tropical dry forest trees in the absence of rain. *Trees*, vol. 16.

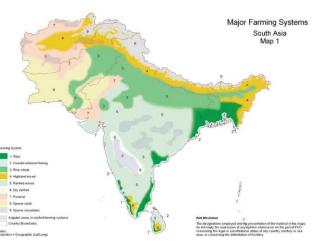
Feldman, Konings, Piles and Entekhabi, 2017.



## Agriculture over Indus Valley and North India



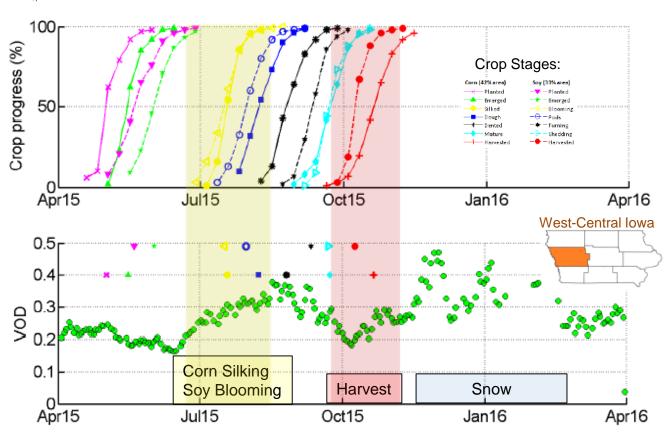






#### Agroecosystems





SMAP measurements are used to map the biomass of crops and vegetation.

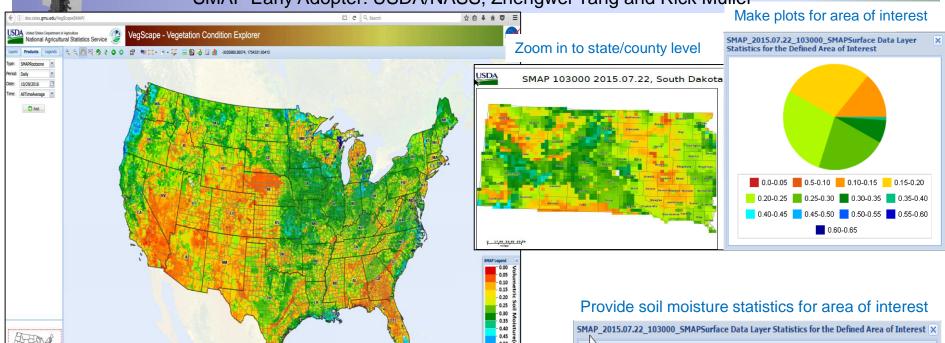
Green biomass in vegetation is proportional to the microwave Vegetation Optical Depth (VOD: green symbols).

The growth phase of corn and soy crops across West-Central lowa correspond to SMAP steady rise in VOD. The harvest period is characterized by sharp drops.

## Improve National Cropland Soil Moisture Monitoring Using SMAP



SMAP Early Adopter: USDA/NASS, Zhengwei Yang and Rick Muller



Using SMAP L4 surface and root-zone soil moisture data, a SMAP prototype on USDA's VegScape, an interactive vegetation condition explorer, illustrates online capabilities to visualize, disseminate, and analyze US cropland soil moisture condition.

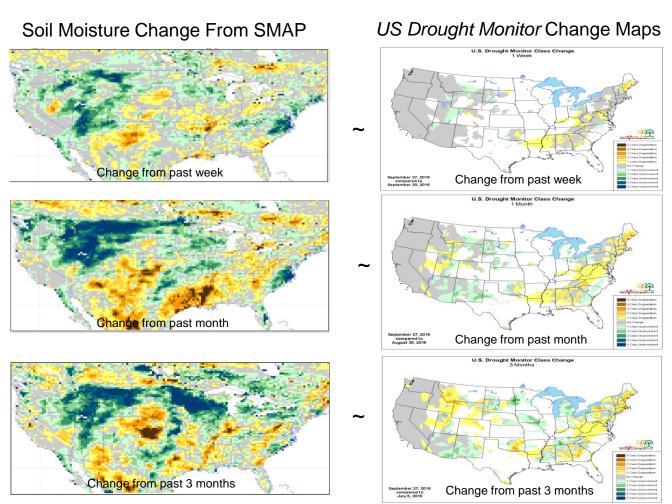
Implementation of VegScape-SMAP will reduce current survey costs and improve the objectivity and robustness of US national soil moisture condition monitoring operations.

Display Crop Area Only				
Value ▲	Category	Acreage	Percentage	
0	0.0-0.05	91128.54	0	
1	0.5-0.10	0	0	
2	0.10-0.15	6766293.81	0.14	
3	0.15-0.20	13281984.14	0.28	
4	0.20-0.25	13122509.2	0.28	
5	0.25-0.30	10388653.12	0.22	
6	0.30-0.35	3235063.03	0.07	
7	0.35-0.40	546771.22	0.01	
8	0.40-0.45	91128.54	0	
9	0.45-0.50	0	0	
10	0.50-0.55	0	0	
11	0.55-0.60	0	0	
Total	13	47523531.6	1.0	



#### Mapping Drought Extent and Recovery



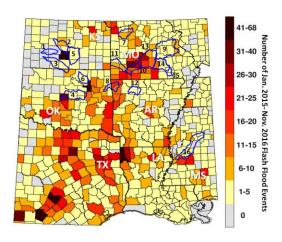


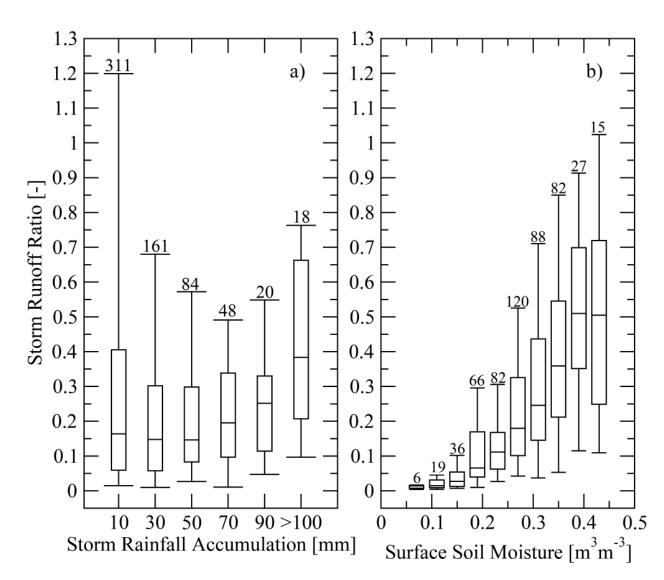
**National Drought Mitigation Center** (NDMC) provides information to State Climatologists who issues maps of drought severity and its recovery status. The SMAP Soil Moisture Change Maps provide NDMC broader coverage and complement the its operational **US** Drought Monitor product.



#### Flood Application



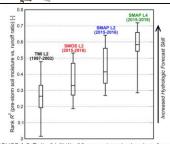




Crow, Chen, Reichle and Liu, 2017: L-band microwave remote sensing and land data assimilation improve the representation of pre-storm soil moisture conditions for hydrologic forecasting, *GRL*.

#### Flood Application





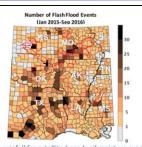
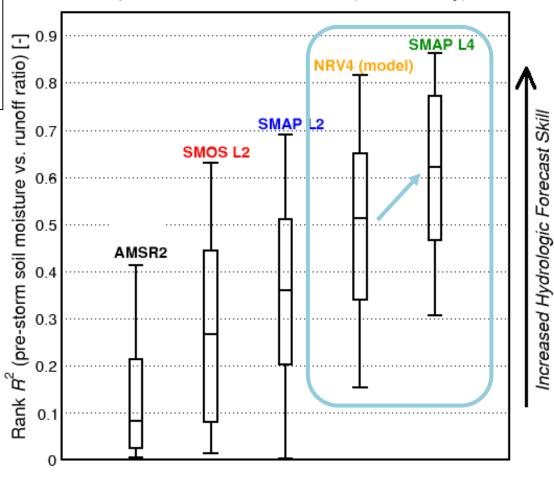


FIGURE 1.2-7: (Left) Skill of forecosting the land surface response to rainfall for satellite-based soil moisture products. Skill is measured as the rank coefficient of determination (R<sup>3</sup>) between pre-storm soil moisture and subsequent storm-scale runoff ratio (i.e., total runoff divided by total rainfall) sampled across 16 basins. The SMAP Level 2 and Level 4 soil moisture products have the highest skill. Relatively lower values for the Tropical Rainfall Measurement Mission Microwave Imager (TMI) and Soil Moisture Ocean Salinity (SMOS) missions reflect skill obtained from earlier sensors. (Right) Boundaries (in red) of the 16 medium-scale (10<sup>3</sup> ±0<sup>4</sup> km²) basins and total number of flash-flood events during January 2015–September 2016 in the south-central U.S.





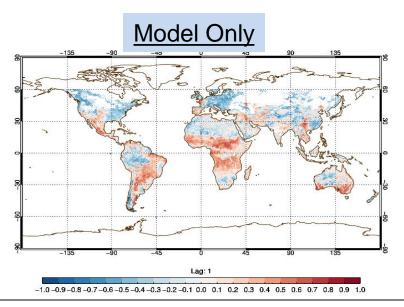
Source: W. Crow et al., 2017.



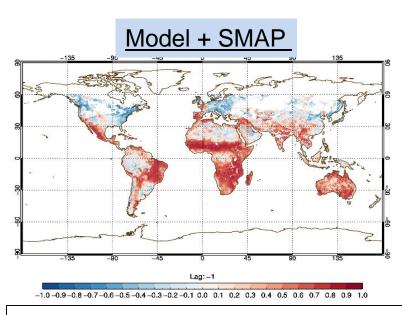
#### USDA FAS (Foreign Agricultural Service) Global Crop Assessment Decision Support System With SMAP Data



Plotted variable = Correlation between *current* monthly soil moisture levels and *future* (+ 1 month) vegetation health (NDVI).



Correlation of <u>current</u> USDA FAS soil moisture product based on water balance modeling



Enhanced correlation observed <u>after</u> the assimilation of SMAP L3 retrievals.

Higher correlation (more red) = Improved early detection of agricultural drought

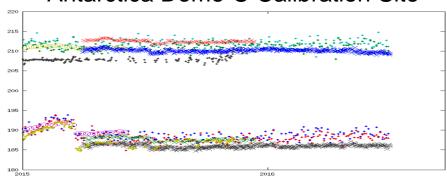
Work supported by a NASA Applied Sciences grant entitled "Enhancing the USDA Global Crop Production Decision Support System with NASA Soil Moisture Active Passive (SMAP) Satellite Observations"

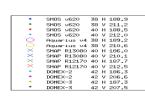


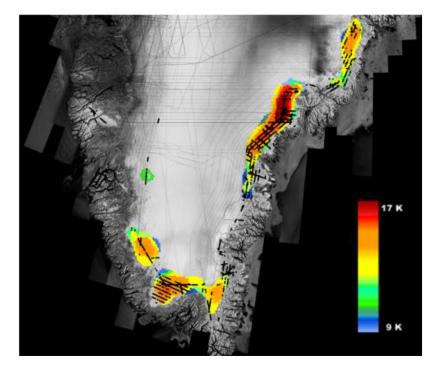
#### Cryosphere



#### Antarctica Dome-C Calibration Site







Miller, J, A. Bringer, K. C. Jezek, J.T. Johnson, R. R. Forster, T. A. Scambos, 2017: Spaceborne observation of Greenland's firn aquifer using L-band microwave radiometry, *Cryosphere* 

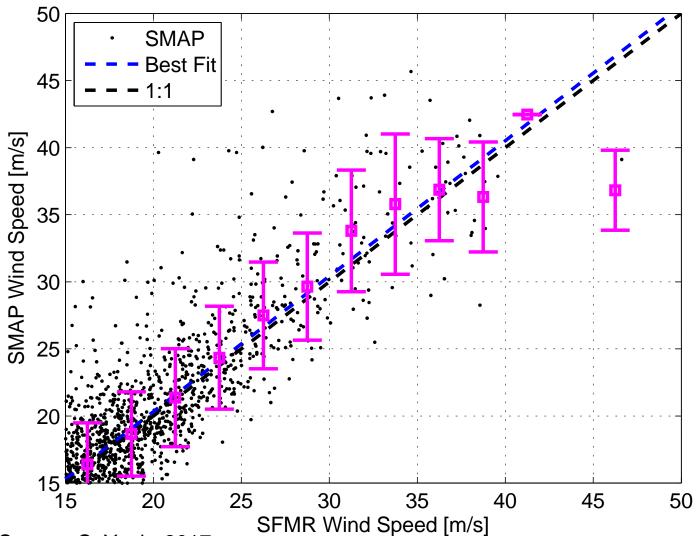
Satellite firn aquifer maps over the Greenland ice sheet retrieved from SMAP using horizontally polarized brightness temperature differences (September average–April average)



#### **SMAP Severe Ocean Wind**



SMAP vs SFMR; Best Fit Slope: 1.01; Corr: 0.83 Mean Error Bar > 20 m/s: 3.15 m/s



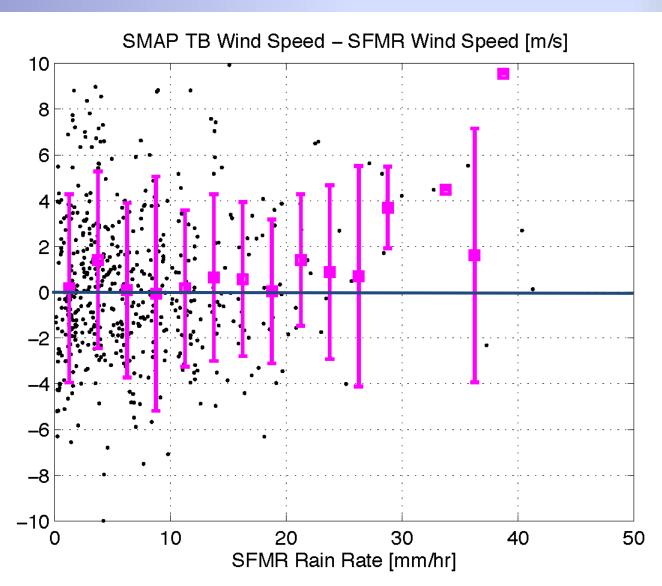
SFMR wind averaged over 40 km along track

Source: S. Yueh, 2017.



#### **SMAP Severe Ocean Wind**





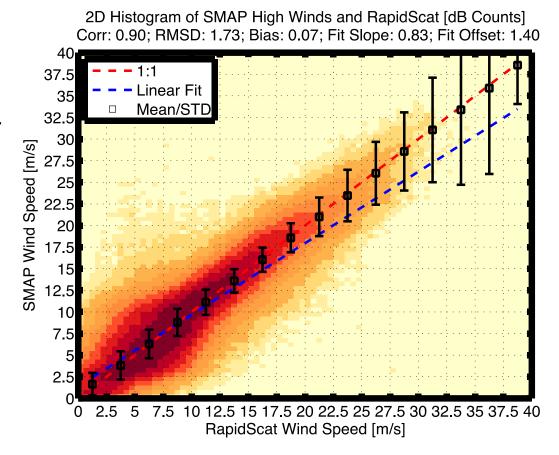
Source: S. Yueh, 2017.



#### SMAP Severe Ocean Wind

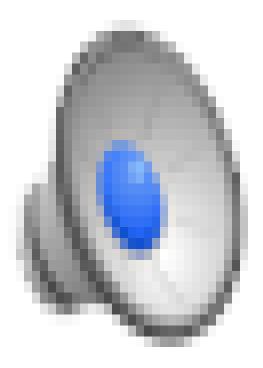


- 13 million matchups of SMAP, RapidScat and Windsat within 90 minute of collocation under rain-free conditions identified by WindSat
- Find very small speed bias up to 40 m/s as compared to Rapidscat.







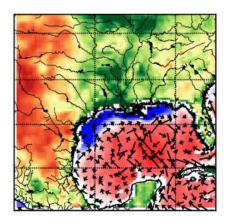


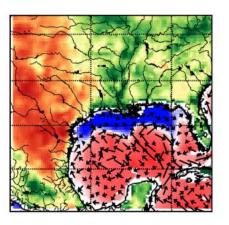


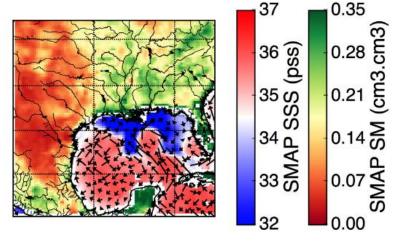
## Linkage of Land and Ocean Branches of Water Cycle



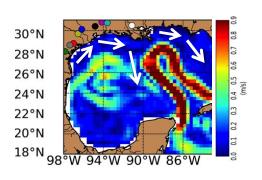
Fournier, S., J. T. Reager, T. Lee, J. Vazquez-Cuervo, C. H. David, and M. M. Gierach, 2016: SMAP observes flooding from land to sea: The Texas event of 2015, *Geophys. Res. Lett.*, 43, 10,338–10,346







SMAP observed an unusual horseshoe-shaped plume of freshwater (dark blue) in the Gulf of Mexico after Texas flooding in May 2015. Louisiana is above the center of the plume, with Florida on the right and the Texas coastline at upper left.





#### Summary



- Exceptional quality global L-band radiometry
- Science uses in characterizing land, terrestrial biosphere and ocean water cycle branches
- Focus on first global characterization of the <u>link</u> between the land branches of the water, energy and carbon cycles: Determines how Earth System models respond to and propagate perturbations in one cycle to another